Peer

A new leptoceratopsid dinosaur from Maastrichtian-aged deposits of the Sustut Basin, northern British Columbia, Canada

Victoria M. Arbour¹ and David C. Evans²

¹ Department of Knowledge, Royal BC Museum, Victoria, BC, Canada

² Department of Natural History, Royal Ontario Museum, Toronto, ON, Canada

ABSTRACT

A partial dinosaur skeleton from the Sustut Basin of northern British Columbia, Canada, previously described as an indeterminate neornithischian, is here reinterpreted as a leptoceratopsid ceratopsian, *Ferrisaurus sustutensis*, gen. et. sp. nov. The skeleton includes parts of the pectoral girdles, left forelimb, left hindlimb, and right pes. It can be distinguished from other named leptoceratopsids based on the proportions of the ulna and pedal phalanges. This is the first unique dinosaur species reported from British Columbia, and can be placed within a reasonably resolved phylogenetic context, with *Ferrisaurus* recovered as more closely related to *Leptoceratops* than *Montanoceratops*. At 68.2–67.2 Ma in age, *Ferrisaurus* falls between, and slightly overlaps with, both *Montanoceratops* and *Leptoceratops*, and represents a western range extension for Laramidian leptoceratopsids.

Subjects Paleontology, Taxonomy

Keywords Dinosauria, Ceratopsia, Cretaceous, Biogeography, Intermontane terrane, Sustut Basin

INTRODUCTION

The dense boreal forest and thrusted, folded rocks of the Canadian Cordillera present a challenging environment in which to search for dinosaurs, compared to the better exposed and more easily accessible outcrops in the badlands of the prairie provinces. Nevertheless, a dinosaur specimen (RBCM P900) consisting of articulated and disarticulated limb and girdle elements was discovered in 1971 in the remote interior mountains of north-central British Columbia (Fig. 1; Arbour & Graves, 2008). These bones were collected by geologist Kenny F. Larsen, who was surveying for uranium along the then in-construction BC Rail line along the Sustut River, and were later donated to Dalhousie University (Halifax, NS) and subsequently accessioned at the Royal British Columbia Museum in Victoria, BC. Arbour & Graves (2008) described this material and identified it as an indeterminate small-bodied, bipedal neornithischian, possibly representing either a pachycephalosaur or a basal ornithopod similar to *Thescelosaurus*. Here, we provide a new interpretation of this material and argue for its assignment to Leptoceratopsidae as a new genus and species. Leptoceratopsids were short-frilled, hornless ceratopsians with a maximum body length of about two-to-three meters, and form the sister group to all other coronosaurian neoceratopsians (*He et al., 2015*). They were present in many Campanian-Maastrictian aged dinosaur assemblages from Asia and

Submitted 24 June 2019 Accepted 20 September 2019 Published 7 November 2019

Corresponding author Victoria M. Arbour, varbour@royalbcmuseum.bc.ca

Academic editor Hans-Dieter Sues

Additional Information and Declarations can be found on page 18

DOI 10.7717/peerj.7926

Copyright 2019 Arbour and Evans

Distributed under Creative Commons CC-BY 4.0

OPEN ACCESS



eer.

Figure 1 RBCM P900, the holotype of *Ferrisaurus sustutensis*, was collected along the BC Rail line near the intersection of Birdflat Creek and the Sustut River in 1971, in the Sustut Basin of northern British Columbia, Canada. Map modified from *Evenchick et al.* (2003). Full-size DOI: 10.7717/peerj.7926/fig-1

North America, but are generally rare in the fossil record (*Ryan et al., 2012; Longrich, 2016*).

RBCM P900 is one of the only vertebrate fossils yet described from the Sustut Basin and as such is significant for understanding the distribution and evolution of dinosaurs in western North America. A 2017 survey of the field area near the confluence of Birdflat Creek and the Sustut River recovered a fragment of the Cretaceous turtle *Basilemys* at a location closely matching Larsen's original field notes, suggesting that RBCM P900 most likely derived from the same outcrop (Fig 1; *Arbour et al., in press*). This work generated new stratigraphic and palynological data that allows the provenance of this important skeleton to be documented in detail for the first time. RBCM P900 is likely from the Tango Creek Formation, rather than the Brothers Peak Formation as originally reported by *Arbour & Graves (2008)*, and the new palynological data suggest that the specimen is late Maastrichtian in age, allowing its morphology and biogeography to be understood in a more detailed temporal context.

METHODS

The electronic version of this article in portable document format will represent a published work according to the International Commission on Zoological Nomenclature (ICZN), and hence the new names contained in the electronic version are effectively published under that Code from the electronic edition alone. This published work and the nomenclatural acts it contains have been registered in ZooBank, the online registration system for the ICZN. The ZooBank Life Science Identifiers (LSIDs) can be resolved and the associated information viewed through any standard web browser by appending the LSID to the prefix http://zoobank.org/. The LSID for this publication is: urn:lsid:zoobank.org: pub:D1C60A34-3632-43AD-BCE0-C93D5E26D1B0. The online version of this work is

archived and available from the following digital repositories: PeerJ, PubMed Central and CLOCKSS. No permits were required for this study and all fossils are permanently accessioned in repositories.

RBCM P900 was compared to ceratopsian, pachycephalosaurid, ornithopodan, and parksosaurid dinosaurs in various collections (Supplemental Information 1) and the literature, and comparative measurements are provided in Supplemental Information 2. Photogrammetric digital models of the specimen (Supplemental Information 3) were created using Agisoft Metashape 1.5.4 using between 50 and 200 digital photos (in RAW format, converted to TIFFs) taken with a Canon Rebel XTi.

We assessed the phylogenetic position of RBCM P900 using the character-taxon matrix for ceratopsians presented by *He et al.* (2015), derived from previous matrices built by *Farke et al.* (2014), *Ryan et al.* (2012), and *Makovicky* (2001). Our matrix includes 34 taxa and 165 characters (Supplemental Informations 1 and 4) and was compiled in Mesquite v3.04 build 725 (*Maddison & Maddison, 2011*). We added three new characters (characters 163–165) based on observations made over the course of this study. We also tested the position of RBCM P900 using the character-taxon matrix presented by *Morschhauser et al.* (2019), with no modifications to the matrix other than the addition of RBCM P900. We performed a cladistic parsimony analysis on both matrices using the Traditional Search option in TNT v1.5 (*Goloboff, Farris & Nixon, 2008*); all characters were treated as unordered and of equal weight, and we used the tree bisection reconnection swapping algorithm with 1,000 replications.

SYSTEMATIC PALAEONTOLOGY

DINOSAURIA Owen, 1842 ORNITHISCHIA Seeley, 1888 NEORNITHSICHIA Cooper, 1985 MARGINOCEPHALIA Sereno, 1986 CERATOPSIA Marsh, 1890 NEOCERATOPSIA Sereno, 1986 CORONOSAURIA Sereno, 1986 LEPTOCERATOPSIDAE Nopcsa, 1923 *FERRISAURUS SUSTUTENSIS* gen. et sp. nov. urn:lsid:zoobank.org:act:A7F4267C-8CC6-49B6-8E52-2C2148929B14

Diagnosis: *Ferrisaurus* can be differentiated from other known leptoceratopsids based on the following unique combination of characters: penultimate pedal phalanges in digits III and IV are equal or subequal in proximodistal length compared to the length of the preceding phalanx, rather than shorter as in all other leptoceratopsids for which these elements are preserved except possibly USNM 13863 (*Cerasinops*); astragalus and tibia coossified, unlike all other leptoceratopsids except for AMNH 5464 (*Montanoceratops*); distal end of ulna broader relative to radius length than in *Leptoceratops;* distal end of ulna medially bowed, unlike the straight ulna of the penecontemporaneous Maastrichtian taxa



Figure 2 Preserved elements of RBCM P900, holotype of *Ferrisaurus sustutensis*, in white (gray represents missing parts of incomplete bones). RBCM P900 includes a partial right coracoid, partial left scapular blade, complete left radius, partial left ulna, partial left tibia, fibula, and coossified astragalus and ?calcaneum, partial left metatarsals I-IV, and digits III (phalanges 2–4) and IV (phalanges 2–5) of the right pes. Full-size DOI: 10.7717/peerj.7926/fig-2

Leptoceratops and *Montanoceratops*, but similar to *Cerasinops* and *Prenoceratops* from the Campanian.

Etymology: "Iron lizard," from Latin *ferrum* (=iron) and Greek *sauros* (=lizard), in reference to the specimen's discovery along a railway line, and *sustutensis* in reference to its provenance near the Sustut River and within the Sustut Basin.

Holotype: RBCM P900, a partial skeleton consisting of a partial right coracoid, fragmentary left scapula, complete left radius, distal portion of the left ulna, associated distal two thirds of the left tibia and fibula and coossified astraglus and ?calcaneum, partial articulated digits III and IV of the right pes, and an unprepared block removed from the posterior surface of the tibia that appears to contain four metatarsals, presumably from the left pes. Previously catalogued as RBCM.EH2006.019.0001 to RBCM.EH2006.019.010 and published under RBCM.EH2006.019 by *Arbour & Graves (2008)*.

Locality: RBCM P900 was discovered near the confluence of Birdflat Creek and the Sustut River in the Sustut Basin (Fig. 1); the bones were found loose in the rubble during construction along the BC Rail line, which has since been abandoned. Fieldwork in the Sustut Basin in 2017 provided strong support for the relocation of the original collection site a few hundred meters from the confluence of the Sustut River and Birdflat Creek (*Arbour et al., in press*); exact GPS coordinates are on file at the Royal BC Museum.

Formation and Age: Tatlatui Member, Tango Creek Formation, Sustut Group. Palynomorphs recovered from the presumed holotype locality included the Maastrichtian marker taxon *Pseudoaquilapollenites bertillonites*, indicating an age of approximately 68.2–67.2 Ma for the site (*Arbour et al., in press*).

LSID: urn:lsid:zoobank.org:act:A7F4267C-8CC6-49B6-8E52-2C2148929B14

DESCRIPTION AND COMPARISON

RBCM P900 includes multiple elements in articulation, including the tibia and fibula, several pedal phalanges, and potentially the metatarsals (Fig. 2). The presence of metatarsals in a block of sediment removed from the posterior face of the distal tibia

suggests that the skeleton may have been fully articulated in situ. The bones do not appear to have suffered from brittle or plastic deformation, but they have been recrystallized, obscuring the original internal bone textures.

We reinterpret RBCM P900 as a leptoceratopsid based on several aspects of the preserved phalanges. The non-ungual phalanges are blockier and more robust in comparison to most orodromines (e.g., Orodromeus MOR 623B), parksosaurids (e.g., Parksosaurus ROM 804), and pachycephalosaurids (e.g., Stegoceras UALVP 2). The dorsal surface of the posterior articular surface in RBCM P900 is more strongly pointed, and overlaps the preceding phalanx more extensively, than in other small ornithischians with ginglymoid phalanges from similar stratigraphic and geographic ranges, such as parksosaurids (e.g., Parksosaurus ROM 804) and pachycephalosaurids (e.g., Stegoceras UALVP 2). Ginglymoid articular surfaces, and narrow, pointed unguals, also exclude identifications of this specimen as a juvenile ceratopsid (e.g., Chasmosaurus UALVP 52613, Currie et al. 2016) or hadrosaurid (e.g., Edmontosaurus annectens, LACM 23504 (Prieto-Márquez, 2014), RAM 7150 (Zheng, Farke & Kim, 2011), Lambeosaurinae indet., TMP 1998.058.0001). The relatively long and robust forelimb compared to the hindlimb, as indicated by the proportions of the radius and tibia, exclude RBCM P900 from being assigned to Thescelosauridae and Pachycephalosauria. The preserved elements of RBCM P900 are comparable in size to large leptoceratopsid specimens like Cerasinops MOR 300 and Leptoceratops CMN 8889.

Pectoral girdle

Arbour & Graves (2008: Figs. 2G and 2H) were unable to identify a thin, gently curved element of RBCM P900, which we reinterpret here as a fragmentary right coracoid (Fig. 3A). Most of the edges are broken, but the angle of the sternal process is complete and part of the anterior edge is complete. The morphology of this bone compares well with the complete coracoids of *Leptoceratops* CMN 8889 (Fig. 3B); the coracoids of most other Laramidian leptoceratopsids are incomplete and cannot be compared with RBCM P900. As in *Leptoceratops*, RBCM P900 had a pronounced, sharply pointed sternal process at the anterior and ventral end of the coracoid. The anterior edge in CMN 8889 (*Leptoceratops*), but without comparable material from other taxa it is difficult to assess whether or not this is within the range of intraspecific variation or a taxonomic difference.

A fragmentary flattened bone was interpreted as a possible rib by *Arbour & Graves* (2008: Figs. 2E and 2F) and is reinterpreted here as part of the left scapula (Figs. 3C and 3D), representing a section near the midpoint of the scapular blade. It has a teardrop-shaped cross section on one side and rapidly narrows to a thin oval cross-section on the other side. The teardrop-shaped outline at one end precludes identification of this element as a rib shaft, and ribs for this individual would have been much smaller and less robust, whereas the proportions are more in line with the scapula of a leptoceratopsid with hindlimb proportions of this size. The ventral edge of the fragment is straight, and the dorsal edge is markedly concave. The scapulae of *Montanoceratops* (MOR 452) and *Prenoceratops* (TCM 2003.1.9 and TCM 2003.1.11; Fig. 3F) are relatively straight along



Figure 3 Pectoral elements of RBCM P900, holotype of *Ferrisaurus sustutensis*, compared to other Laramidian leptoceratopsids. (A) Fragmentary right coracoid of RBCM P900 in lateral view, compared to (B) complete right scapulocoracoid of CMN 8889, *Leptoceratops gracilis*, lateral view centered on coracoid with scapula in oblique view. Fragmentary left scapular blade of RBCM P900 in (C) lateral and (D) medial view, compared to (E) left scapula of MOR 300, *Cerasinops hodgskissi* in medial view, and (F) left scapula of TCM 2003.1.9, *Prenoceratops pieganensis* in lateral view. Abbreviations: sp, sternal process. Full-size DOI: 10.7717/peerj.7926/fig-3

their dorsal lengths, whereas the scapulae of *Cerasinops* (MOR 300, Fig. 3E) and *Leptoceratops* (CMN 8889) are more concave dorsally in lateral view.

Forelimb

We agree with the identification of the radius by *Arbour & Graves (2008*: Figs. 2C and 2D). The radius is a relatively simple rod-shaped bone with gently expanded proximal and distal ends and a shaft that is triangular in cross section (Figs. 4A–4D; Table 1). Overall, the



Figure 4 Radius of RBCM P900, holotype of Ferrisaurus sustutensis, compared to other Laramidian leptoceratopsids. RBCM P900, Ferrisaurus sustutensis, left radius in (A) lateral, (B) medial, (C) proximal, and (D) distal view. (E) CMN 8889, Leptoceratops gracilis, left radius in lateral view. (F) MOR 300, Cerasinops hodgskissi, ?left radius in ?lateral view. Abbreviations: tb, tubercle.

Full-size DOI: 10.7717/peerj.7926/fig-4

Table 1 Selected measurements of forelimb and hindlimb elements in leptoceratopsids (mm).								
Taxon	Specimen	Radius	Ulna		Tibia		Fibula	Measurement source
		Length	Length	Distal width	Length	Distal width	Length	
Ferrisaurus sustutensis	RBCM P900	135.0		38.2		90.1		Direct measurement
Cerasinops hodgskissi	MOR 300 R		201.4	>32.6	363.0	~86.3	337.0	Direct measurement
	MOR 300 L					95.0		Direct measurement
	USNM 13863				200	62		Brown & Schlaikjer (1942)
Ischioceratops zhuchengensis	ZCFM V0016				329			He et al. (2015)
Leptoceratops gracilis	AMNH 5205	167	224			117		Sternberg (1951)
	CMN 8887	115			240			Sternberg (1951)
	CMN 8888	137			290			Sternberg (1951)
	CMN 8889 L	160.5	202.5	35.7	323.0	87.9	293.0	Direct measurement
	PU 18133				~385	~78		Ostrom (1978)
Montanoceratops cerorhynchus	AMNH 5464				355	102		Brown & Schlaikjer (1942)
	MOR 542			28.6	249.2	50.6	235.9	Direct measurement
Prenoceratops pieganensis	TCM 2003.1.8		143.3	19.1				Direct measurement

. . . .



Figure 5 Ulna of RBCM P900, holotype of *Ferrisaurus sustutensis*, compared to other Laramidian **leptoceratopsids.** RBCM P900, *Ferrisaurus sustutensis*, left ulna in (A) medial and (B) distal view. (C) CMN 8889, *Leptoceratops gracilis*, left ulna in medial view. (D) MOR 300, *Cerasinops hodgskissi*, right ulna in medial view. (E) TCM 2003.1.8, *Prenoceratops pieganensis*, right ulna in medial view. (F) MOR 452, *Montanoceratops cerorhynchus*, right ulna in medial view. (G) RBCM P900, *Ferrisaurus* left ulna in posterior view; arrow indicates medial bend to distal ulna. (H) TCM 2003.1.8, *Prenoceratops* right ulna in anterior view. (I) MOR 452, *Montanoceratops* right ulna in anterior view.

Full-size 🖾 DOI: 10.7717/peerj.7926/fig-5

radius of RBCM P900 is very similar to that of *Leptoceratops* (*Brown, 1914*; Fig. 4E), and it differs only in subtle aspects. The proximal end in RBCM P900 is less cup-shaped compared to *Leptoceratops* (CMN 8889), and the shaft lacks the prominent protuberance present near the midpoint in *Leptoceratops* (AMNH 5205; *Brown, 1914*), although a light distal tuberosity is present as in AMNH 5205. The preserved radii of *Cerasinops* (MOR 300; Fig. 4F) lack distal and proximal ends, but preserve straight shafts lacking any bulges or tuberosities.

We reinterpret the bone previously identified by *Arbour & Graves* (2008: Fig. 2) as the proximal half of a humerus as a partial right ulna including the distal end (Figs. 5A and 5B; Table 1). The absence of a prominent deltopectoral crest or rounded humeral head is inconsistent with its identification as a humerus. The ulna is incomplete proximally, but the shaft is expanded toward the broken proximal end. Based on the proportions of the radius length to ulna length in *Leptoceratops, Montanoceratops*, and to a lesser extent *Cerasinops* (Supplemental Information 2) where the radius is 75–80% of the length of the ulna, the ulna of RBCM P900 may have been 170–180 mm in total length. The proximal expansion of the ulna occurs approximately 100 mm from the base of this element in RBCM P900, compared to about 120 mm in CMN 8889 (*Leptoceratops*; 67% of the length), and 125 mm in MOR 300 (*Cerasinops*; 62% of the length). Extrapolating a total length for the ulna of RBCM P900 based on these proportions yields a total length of ~150–170 mm. Comparing the width of the distal ulna to the length of the radius, the ulna of RBCM

P900 was proportionately wider compared to other leptoceratopsids (Fig. 5; Supplemental Information 2), giving it a stouter appearance.

The ulna shaft is a flattened oval in cross-section, and the distal end is flat and only moderately expanded. A diagnostic character for *Cerasinops* proposed by *Chinnery* & *Horner* (2007) is the strong medial bend of the distal part of the ulna. The distal ulna of RBCM P900 is also medially deflected (Fig. 5G), with the posterior edge more strongly curved than the anterior edge. The postcrania of the bonebed material of *Prenoceratops* was not previously described by *Chinnery* (2004), but examination of TCM 2003.1.8, a right ulna (Fig. 5H), indicates that *Prenoceratops* also had a medial bend to the distal ulna. The ulna is straight in this region in *Leptoceratops* (CMN 8889) and *Montanoceratops* (MOR 542; Fig. 5I).

Hindlimb

Approximately the distal two thirds of the right tibia and fibula are preserved, with the tibia and fibula in articulation (Figs. 6A–6E; Table 1). Using more complete specimens of similar size as a guide (Supplemental Information 2), we estimate that the tibia in RBCM P900 was likely between 310 and 330 m in length originally. The astragalus and possibly the calcaneum are coossified to the tibia (Fig. 6C) but the boundaries between these elements are difficult to discern. The tibia and astragalus are not coossified in *Leptoceratops* (CMN 8889; Figs. 6H and 6I), *Cerasinops* (MOR 300; Figs. 6J–6L) or *Montanoceratops* (MOR 542) and in these specimens the boundary between these elements is clearly discernible. *Makovicky (2010)* notes that the astragalus is partly coossified with the tibia in *Montanoceratops* (AMNH 5465). It is unclear whether this is an ontogenetic phenomenon, and if it is phylogenetically significant.

In medial and lateral views (Figs. 6A and 6E) the tibia of RBCM P900 has a pronounced distal curvature that was not observed in any other leptoceratopsid specimens and which does not seem to represent taphonomic deformation, based on the absence of crushing or fractures on the tibia. In distal view (Fig. 6B), the lateral and medial malleoli are offset at a distinct angle, giving the distal face of the tibia/astragalus a triangular cross section; RBCM P900 has a more pronounced edge marking the confluence between the malleoli compared to the condition in Leptoceratops (CMN 8889), Cerasinops (MOR 300), or Montanoceratops (MOR 542). The tibia of RBCM P900 is straight-sided in anterior and posterior view and tapers toward the midpoint in anterior or posterior view, similar to the condition in Leptoceratops (CMN 8889; Figs. 6H and 6I), and Montanoceratops (MOR 542), and unlike the strongly kinked morphology observed in Cerasinops (MOR 300; Figs. 6J-6L). The tibia narrows significantly along the shaft and has an oval cross section at its broken proximal end. The fibula is narrow, with an oval cross section. A portion of matrix removed from the anterior side of the distal tibia contains what appear to be the remains of four metatarsals in cross section (Figs. 6F and 6G), but little can be said about their morphology without further preparation.

RBCM P900 preserves a large number of pedal phalanges: III-2, III-3, and III-4, and IV-2, IV-3, IV-4, and IV-5 (Figs. 7A–7C; Table 2). Pedal digit III was preserved in articulation on a piece of matrix (Figs. 7A and 7B); digit IV includes IV-2 and IV-3

PeerJ



Figure 6 Tibia of RBCM P900, holotype of *Ferrisaurus sustutensis*, compared to other Laramidian leptoceratopsids. RBCM P900, *Ferrisaurus* left tibia in (A) medial, (B) and (C) posterior, (D) anterior, and (E) lateral views, and (F) and (G) block removed from anterior face of tibia containing four partial metatarsals. The dashed line in (C) delineates the possible boundary of the astragalus/calcaneum on the tibia, and the dashed lines in (E) indicate the preserved metatarsals in cross-section. CMN 8889, *Leptoceratops gracilis* left tibia in (H) posterior and (I) anterior view. MOR 300, *Cerasinops hodgskissi* right tibia in (J) anterior and (K) posterior views, and (L) left tibia in posterior view. Abbreviations: as, astraglus; ca, calcaneum; fib, fibula; ma, matrix; mt, metatarsal.



Figure 7 Pedal elements of RBCM P900, holotype of *Ferrisaurus sustutensis*, compared to other Laramidian small-bodied ornithischians. RBCM P900, *Ferrisaurus*, left digit III in (A) medial and (B) lateral views, and (C) left digit IV in lateral view. (D) MOR 542, *Montanoceratops cerorhynchus*, right digit IV in lateral view. Illustrations of (E) RBCM P900, *Ferrisaurus* sustutensis, (F) CMN 8889, *Leptoceratops gracilis*, (G) MOR 300, *Cerasinops hodgskissi*, (H) MOR 542, *Montanoceratops cerorhynchus*, and (I) ROM 804, *Parksosaurus warreni*, in dorsal view. Full-size DOI: 10.7717/peerj.7926/fig-7

> preserved in articulation and IV-4 and IV-5 can be "snapped" back into articulation based on the presence of some remaining matrix on these elements (Fig. 7C). The non-ungual phalanges are somewhat longer than wide, but blocky rather than elongate, and ginglymoid, distinguishing them from similarly-sized small-bodied ornithischians such as *Parksosaurus* (Fig. 7I). The distinctly ginglymoid nature of the interphalangeal joints is distinct from the non-ginglymoid pedal phalangeal joints in Hadrosauridae (*Zheng, Farke & Kim, 2011*).

> In *Leptoceratops* (CMN 8889, CMN 8887), *Cerasinops* (MOR 300), and *Montanoceratops* (MOR 542, Fig. 7D) the penultimate pedal phalanx of each major digit is markedly shorter in length compared to the preceding phalanx (~75–90% the length of the preceding phalanx); in RBCM P900 the penultimate and preceding phalanx on digits III and IV are similar in size, with the penultimate phalanx actually being slightly longer than the preceding phalanx (Table 2; Supplemental Information 2). *Leptoceratops* (AMNH 5205; *Brown, 1914*) and *Cerasinops* (USNM 13863; *Gilmore, 1939*;

Table 2 Lengths of phalanges from pedal digits III and IV in leptoceratopsids (mm).									
Taxon	Specimen	III		IV				Measurement source	
		2	3	4	2	3	4	5	
Ferrisaurus sustutensis	RBCM P900	28.1	28.3	40.7	24.4	21.1	22.3	29.3	Direct measurement
Cerasinops hodgskissi	MOR 300 R	?34.1	?27.9			?25.4			Direct measurement
	MOR 300 L	36.5	29.8	44.9	33.5	27.2	20.1	>31.1	Direct measurement
	USNM 13863	27.5	29.5	41	21	21	18.5		Brown & Schlaikjer (1942)
Leptoceratops gracilis	CMN 8887	21.3	16.5	32.2	18.4	14.4	13.7	24.9	Direct measurement from cast
	CMN 8889 R	31.9	29.6	51.0	29.4	25.1	22.0	44.1	Direct measurement
	PU 18133	40	30		32	~26	~20		Ostrom (1978)
Montanoceratops cerorhynchus	AMNH 5464	33		68					Brown & Schlaikjer (1942)
	MOR 542	28.4	25.1	~29.0	20.4	21.0	19.1	34.0	Direct measurement

Chinnery & Horner, 2007) are both illustrated with penultimate phalanges subequal in length to the preceding phalanx, but these are both illustrated as line drawings rather than photographs, measurements were not provided by the authors, the digits in AMNH 5205 were not part of an articulated pes, and neither of these specimens were measured for this study. As such, it is unclear if the illustrations accurately reflect the actual morphology of the pedal digits in these two specimens. The figured pes of *Udanoceratops* PIN 4046/11 (*Tereschenko, 2008*) appears to show penultimate phalanges subequal in length to the preceding phalanx in digits II-IV, but measurements were not provided, and no rationale was provided for why this specimen is referred to *Udanoceratops* rather than *Protoceratops*. RBCM P900 can, however, be differentiated from *Udanoceratops* by the morphology of the pedal unguals, if PIN 4046/11 (*Tereschenko, 2008*) is referable to *Udanoceratops* rather than *Protoceratops* rather than *Protoceratops*.

The unguals of RBCM P900 are long and narrow, with a gently curved ventral surface (Figs. 7A–7C and 7E), differing from the broad, hoof-shaped unguals of ceratopsids or the wide triangular unguals of protoceratopsids (*Sternberg, 1951*). Their overall shape is similar to the unguals of most other leptoceratopsids, with the possible exception of *Udanoceratops* based on specimen PIN 4046/11 where the proximal articular surface of the ungual is much wider than the distal articular surface of the penultimate phalanx (*Tereschenko, 2008*). Lateral grooves on the unguals of RBCM P900 are shallow. The unguals of *Leptoceratops* specimen CMN 8889 have a longitudinal furrow on the ventral surface, but these are absent in the smaller *Leptoceratops* specimen CMN 8887, and ventral furrows were not observed on any other leptoceratopsid unguals examined for this study. No ventral furrows are present on the unguals of RBCM P900. The unguals of RBCM P900 appear slightly deeper in lateral view compared to other leptoceratopsids, but it is unclear how much this is influenced by taphonomic factors (e.g., *Montanoceratops* MOR 542 is substantially smaller than RBCM P900).

RESULTS OF THE PHYLOGENETIC ANALYSES

The phylogenetic analysis of the *He et al. (2015)* modified matrix recovered seven most parsimonious trees, each with a tree length of 328, a consistency index of 0.60, a retention





index of 0.80, and a best tree-bisection reconnection score of 326 (Fig. 8A). The strict consensus tree (Fig. 8) is nearly identical to that presented by *He et al. (2015)*, with a basal grade of small-bodied ceratopsians and two derived clades, Coronosauria and Leptoceratopsidae. Within Leptoceratopsidae, the recovered relationships are similar to those found by *He et al. (2015)*, with *Asiaceratops* and *Cerasinops* recovered as successive sister taxa to all other leptoceratopsids, *Montanoceratops* and *Ischioceratops* as sister taxa, and *Prenoceratops* as the sister taxon to an unresolved clade of the six remaining leptoceratopsids, including *Ferrisaurus*. Within this clade, *Ferrisaurus* has an unresolved relationship with the North American taxa *Leptoceratops*, *Gryphoceratops* and *Unescoceratops* and the Asian taxa *Udanoceratops* and *Zhuchengceratops*. Poor resolution of this group is most likely because of the low number of characters that could be coded for *Ferrisaurus*. In two of the seven trees, *Ferrisaurus* and *Udanoceratops* were sister taxa; the position of *Ferrisaurus* differs in the other five trees. Moving *Ferrisaurus* basally in the tree, outside of Coronosauria + Leptoceratopsidae, increases the tree length to 329, and moving *Ferrisaurus* into Ceratopsidae increases the tree length to 331.

<u>'eer</u>

The analysis of the *Morschhauser et al. (2019)* unmodified matrix recovered 1,110 most parsimonious trees, each with a tree length of 694, a consistency index of 0.45, a retention index of 0.67, and a best tree-bisection reconnection score of 688. *Morschhauser et al.'s (2019)* strict consensus tree shows a poorly resolved sister clade to Coronosauria consisting of taxa typically recovered as leptoceratopsids in other analyses plus *Koreaceratops* and *Helioceratops* (Fig. 8B). The addition of *Ferrisaurus* to this matrix collapses this clade, and *Ferrisaurus* is recovered in an unresolved polytomy of leptoceratopsids plus *Aquilops, Archaeoceratops, Auroraceratops, Helioceratops*, and *Koreaceratops*, outside of Coronosauria. In 63% of the trees, *Ferrisaurus* is recovered as a leptoceratops, *Leptoceratops, Montanoceratops, Prenoceratops, Udanoceratops, Zhuchengceratops*, and *Gryphoceratops + Unescoceratops*, with *Helioceratops* as the outgroup.

DISCUSSION

The fact that RBCM P900, the first dinosaur specimen recovered from the Sustut Basin, is a leptoceratopsid rather than one of the more commonly encountered groups in many coeval formations in western North America, such as hadrosaurs, ceratopsids, or tyrannosaurs, is surprising, especially given well-documented preservational biases against small-bodied dinosaurs in more fossiliferous areas (*Brown et al., 2013a, 2013b; Evans et al., 2013*). Most leptoceratopsid taxa are distinguished on the basis of cranial morphology, especially aspects of the lower jaw anatomy (*Ryan et al., 2012*). However, excellent postcranial material is known for many taxa, making it possible to identify diagnostic features in RBCM P900 despite the absence of cranial material for this specimen. *Leptoceratops, Montanoceratops,* and *Cerasinops* are all known from multiple partial or complete skeletons (*Chinnery & Weishampel, 1998; Chinnery & Horner, 2007; Ostrom, 1978; Brown & Schlaikjer, 1942; Sternberg, 1951; Brown, 1914*), and *Prenoceratops* specimens described by *Chinnery (2004)* come from a single mixed bonebed from which multiple composite skeletons have been assembled.

Digit proportions have been used to distinguish caenagnathids (*Zanno & Sampson, 2005*), oviraptorids (*Longrich, Currie & Zhi-Ming, 2010*), and ornithomimids (*Kobayashi & Barsbold, 2006*) at low taxonomic levels, and we show that they can also be used to distinguish among leptoceratopsids. In all specimens preserving partial or complete articulated pedes that we were able to personally observe and measure, the penultimate phalanx (preceding the ungual) for each major digit is shorter in length than the immediately preceding phalanx. In other words, pedal phalanx length decreases distally in the digit, except for the unguals (Fig. 7). In *Ferrisaurus*, the penultimate phalanx is subequal in length to the preceding phalanx in digits 3 and 4, and phalanx length does not decrease distally within each pedal digit. This appears to be unique to *Ferrisaurus* within leptoceratopsids with two possible exceptions. This morphology may be present in a referred specimen of *Udanoceratops* (PIN 4046/11, *Tereschenko, 2008*), although it is not clear that this specimen is not referable to *Protoceratops*, and measurements were not provided. *Gilmore (1939)* published measurements for USNM 13863 (*Cerasinops*) and noted the length of III-2 as 27.5 and III-3 as 29.5 mm; although we have not had the opportunity to observe this specimen in

person, a two mm length increase between III-2 and III-3 is far outside the range of variation we observed in leptoceratopsids over the course of this study (Table 2; Supplemental Information 2), but is within the range of variation of a decrease in length between III-2 and III-3. Additionally, phalanges in digit IV show the more typical reduction in length distally. We suggest it is possible that III-2 and III-3 in USNM 13863 were at some point transposed in their positions, despite the pes being reported as articulated at the time of collection by *Gilmore (1939)*. Longer penultimate phalanges may also be present in more basal ceratopsian taxa such as *Archaeoceratops (Dodson, 2003)*, although phalangeal measurements were not provided in the descriptions of this taxon; phalangeal length decreases in *Yinlong* as for leptoceratopsids except *Ferrisaurus (Han et al., 2018)*. Overall, the observed pattern for leptoceratopsids appears to be a marked decrease in non-ungual phalangeal length in each pedal digit, with the exception of *Ferrisaurus*.

The astragalus and tibia in RBCM P900 are coossified (Fig. 6), an unusual condition among leptoceratopsids that is otherwise reported in only one specimen of *Montanoceratops* (AMNH 5465). Coossification of the astragalus and tibia could indicate advanced skeletal maturity in RBCM P900, but this specimen is smaller than specimens in which the tibia and astragalus remain separate (e.g., *Leptoceratops* CMN 8889, *Montanoceratops* AMNH 5205), suggesting that size alone does not explain the differences in coossification patterns in leptoceratopsids. It is unclear what the ontogenetic significance of this coossification represents in *Ferrisaurus*. Fusion of the ankle (distal tibia and fibula) has been proposed as a diagnostic character of the small bodied thescelosaurid *Albertadromeus syntarsus* from the Campanian of Alberta (*Brown et al., 2013b*), and a distinctive feature of some coelophysoids such as "*Syntarsus*"/?*Coelophysis kayentakatae* (*Rowe, 1989*), mature derived ceratosaurs, such as *Cryolophosaurus* (*Smith et al., 2007*), and mature ankylosaurs (*Coombs, 1971*) and ceratopsids (*Sues & Averianov, 2009*).

Ferrisaurus shares with Cerasinops a medially bent distal ulna (originally proposed as a diagnostic character for Cerasinops by Chinnery & Horner, 2007), a feature that is also present in Prenoceratops (TCM 2003.1.8). This feature is not present in the Maastrichtianaged leptoceratopsids Montanoceratops and Leptoceratops, which are closest in geological age to Ferrisaurus. Chinnery & Horner (2007) suggested that the medial deflection of the ulna in *Cerasinops*, as well as the proportions and histology of the limb elements, may indicate that Cerasinops was primarily bipedal rather than quadrupedal. Although limb proportions are more difficult to determine in *Ferrisaurus*, if the complete tibia was between 310 and 330 mm (estimated based on more complete tibiae in Leptoceratops and Montanoceratops, Supplemental Information 2), then the radius of *Ferrisaurus* would have been no more than 40-43% of the length of the tibia. This is less than other comparable leptoceratopsids: the radius is 50% the length of the tibia in Leptoceratops CMN 8889, 48% in Leptoceratops AMNH 5205, and 47% in Leptoceratops CMN 8888, and much more than 45% in the incomplete radii of Cerasinops MOR 300. Ferrisaurus thus may have had a more robust distal ulna (Fig. 5), but a shorter forelimb overall compared to *Cerasinops*, suggesting that it too may have been at least facultatively bipedal. Alternately, the robusticity of the ulna may be related to another aspect of its ecology, such as digging, which has been suggested in the orodromine Oryctodromeus (Fearon & Varricchio, 2015) and Protoceratops (Longrich, 2010).

Ferrisaurus was recovered as a leptoceratopsid using the modified *He et al.* (2015) character matrix, and as a non-coronosaurian neoceratopsian in the *Morschhauser et al.* (2019) matrix. Although the precise relationships of *Ferrisaurus* are unresolved using the *He et al.* (2015) matrix, we found it to be more closely related to *Leptoceratops* than *Montanoceratops* (Fig. 8). Despite their stratigraphic and geographic proximity, *Leptoceratops* and *Montanoceratops* are not recovered as close relatives in recent phylogenetic analyses in this analysis or by *He et al.* (2015) and preceding versions of that matrix. *Montanoceratops* occupies a relatively basal position within Leptoceratopsidae (*Makovicky, 2010; Ryan et al., 2012; Farke et al., 2014; He et al., 2015). Leptoceratops* typically occupies a more derived position and has been recovered as the sister taxon to the Asian *Udanoceratops* (*He et al., 2015*). *Ferrisaurus* was thus recovered in a more derived position within Leptoceratops.

Stratigraphic and palaeobiogeographic implications

Leptoceratopsids are known from the Santonian through Maastrichtian of Laramidia (Ryan et al., 2012), and the Campanian-Maastrichtian of Mongolia and China (He et al., 2015); fragmentary putative leptoceratopsids have also been reported from the Cenomanian of Uzbekistan (Nessov, Kaznyshkina & Cherepanov, 1989), the ?Coniacian-Santonian of Belgium (Godefroit & Lambert, 2007; Longrich, 2016), the Campanian of North Carolina (Longrich, 2016), and the Campanian of Sweden (Lindgren et al., 2007). The ancestor of the leptoceratopsid lineage most likely originated in Asia (*Chinnery*-Allgeier & Kirkland, 2010), but multiple exchanges across Beringia from Asia to North America (and vice versa) may have occurred. Gryphoceratops, the oldest taxon, derives from the Deadhorse Coulee Member of the Milk River Formation, with a minimum age of about 83.7 Ma (Ryan et al., 2012). Campanian Laramidian taxa include Cerasinops from the lower Two Medicine Formation, Prenoceratops from the upper Two Medicine Formation of Montana and the Oldman Formation of Alberta, and Unescoceratops from the lower Dinosaur Park Formation (Chinnery, 2004; Chinnery & Horner, 2007; Ryan et al., 2012). Only two genera are known from the Maastrichtian of Laramidia: Montanoceratops from the St Mary River and Horseshoe Canyon formations (Brown & Schlaikjer, 1942; Chinnery & Weishampel, 1998; Makovicky, 2001), and Leptoceratops from the Scollard and Hell Creek formations (Sternberg, 1951; Ott, 2007) and the Pinyin Conglomerate (McKenna & Love, 1970). RBCM P900 was most likely collected from approximately 68.2-67.2 Ma sediments of the Tatlatui Member of the Tango Creek Formation, based on a recent field reassessment of its original collection locality and palynomorphs recovered from that site (Arbour et al., in press). This places it between the stratigraphic ranges for Montanoceratops (71.939-68 Ma) and Leptoceratops (66.97-66 Ma), and slightly overlapping with the known range of Montanoceratops (Fowler, 2017).

Stratigraphically, *Montanoceratops* and *Leptoceratops* are the most likely taxa to which RBCM P900 could be referred, but multiple anatomical features distinguish RBCM P900 from both *Leptoceratops* and *Montanoceratops*, including the proportions of the pedal digits, the proportions of the ulna, and the medially bowed morphology of the distal ulna.

RBCM P900 is also unlikely to represent an individual of *Cerasinops* or *Prenoceratops*; it can be distinguished from *Cerasinops* based on the proportions of the pedal digits, and from both *Cerasinops* and *Prenoceratops* based on the proportions of the ulna. These morphological differences are reinforced by the stratigraphic position of *Ferrisaurus* relative to *Cerasinops* and *Prenoceratops* (latest Maastrichtian, vs. middle to Upper Campanian; *Chinnery & Horner*, 2007; *Chinnery*, 2004), given that no other dinosaur species with temporally well-resolved specimens spans the middle Campanian to latest Maastrichtian elsewhere in Laramidia (*Eberth et al.*, 2013; *Fowler*, 2017). An enigmatic specimen, TMP 1982.11.1, from the Maastrichtian Willow Creek Formation (*Miyashita, Currie & Chinnery-Allgeier*, 2010) has been referred to *Montanoceratops* by several authors (*Ryan & Currie*, 1998), but was considered neither a representative of *Montanoceratops*, *Leptoceratops*, or *Cerasinops* by *Makovicky* (2010). Several additional as-yet undescribed specimens in the collections of the TMP (*Tanke*, 2007) may represent examples of either *Montanoceratops*, *Leptoceratops*, or *Ferrisaurus* and their description may help clarify the differences between these three taxa or provide new anatomical information for *Ferrisaurus*.

Leptoceratopsids are uncommon components of the dinosaurian faunas of Laramidia: even in the well-sampled Dinosaur Park Formation of Alberta only a handful of leptoceratopsid specimens are known (*Tanke, 2007*). *Ryan & Evans (2005)* hypothesized that leptoceratopsids may have avoided the wet coastal environments favored by ceratopsids. *Leptoceratops* appears to be present primarily in piedmont and alluvial plain palaeoenvironments and is largely absent in coastal plain settings (*Lehman, 1987*, although see *Ott, 2007*). The Tatlatui Member of the Tango Creek Formation represents an alluvial plain palaeoenvironment (*Bustin & McKenzie, 1989*), consistent with the palaeoenvironmental association documented for other Maastrichtian leptoceratopsids. Interestingly, the intermontane basin occurrence of *Ferrisaurus* also supports one hypothesis outlined by *Lehman (1987, 2001*), that leptoceratopsids, along with a few other large-bodied herbivorous taxa, were inhabitants of Cordilleran highlands and adjacent piedmonts, which, in part, explains their rarity in the fossil record.

Although today the holotype locality for *Ferrisaurus* is found at approximately 56°N, the unusual and complex translational history of the Intermontane Superterrane means its palaeolatitude may have lain as much as 1,600 km to the south of its current position with respect to cratonic North America, and may have had approximately the same palaeolatitude (~48°N) as the southern border of Oregon and Idaho (*Enkin et al., 2003*; *Van Hinsbergen et al., 2015*). Despite its current apparent northern latitude, the holotype of *Ferrisaurus* may actually represent one of the southernmost occurrences of Leptoceratopsidae in western North America, and at minimum would have been within the currently known latitudinal range of Laramidian leptoceratopsids. Regardless, RBCM P900 represents a western range extension for Laramidian leptoceratopsids, and a unique occurrence within a restricted intermontane basin palaeoenvironment. The identification of RBCM P900 as a unique leptoceratopsid distinct from other known Laramidian taxa supports previous conclusions by *Makovicky (2010)* and *Ryan et al. (2012)* that Leptoceratopsidae was a diverse but currently poorly sampled lineage of Late Cretaceous ceratopsians.

CONCLUSIONS

RBCM P900, previously identified as an indeterminate bipedal neornithischian by *Arbour* & *Graves* (2008), instead represents the partial skeleton of a leptoceratopsid ceratopsian similar in size to large specimens of *Leptoceratops* and *Cerasinops*. Although fragmentary, this specimen can be differentiated from other leptoceratopsids based on the proportions and morphology of the ulna and pedal digits, and is designated the holotype of the new taxon *Ferrisaurus sustutensis*. RBCM P900 was collected from the Sustut Group of the southern Sustut Basin, a large but relatively unexplored terrestrial Cretaceous basin in northern British Columbia, Canada. Its recognition as a distinct species of a generally rare group of small-bodied dinosaurs highlights the potential for future discoveries of unique dinosaur biodiversity within the intermontane basins of the western side of the North American Cordillera.

INSTITUTIONAL ABBREVIATIONS

LACM	Los Angeles County Museum
MOR	Museum of the Rockies, Bozeman, Montana, USA
RBCM	Royal BC Museum, Victoria, British Columbia, Canada; Raymond M. Alf
	Museum of Paleontology, Claremont, California, USA
ROM	Royal Ontario Museum, Toronto, Ontario, Canada
ТМР	Royal Tyrrell Museum of Palaeontology, Drumheller, Alberta, Canada
UALVP	University of Alberta, Edmonton, Alberta, Canada
CMN	Canadian Museum of Nature, Ottawa, Ontario, Canada.

ACKNOWLEDGEMENTS

The RBCM P900 field locality is located on the unceded traditional territory of the Gitxsan peoples. MOR 300 was collected from the Wilson Hodgkiss Ranch, and MOR 542 was collected from private land deeded from the Blackfeet Nation. Many thanks to Dallas Evans (Children's Museum of Indianapolis), Jordan Mallon and Kieran Shepherd (Canadian Museum of Nature), Amy Atwater, Scott Williams, and John Scannella (Museum of the Rockies), Brandon Strilisky and Caleb Brown (Royal Tyrrell Museum of Palaeontology), and Carl Mehling (American Museum of Nature History) for access to specimens in their collections. Peter Makovicky shared photographs of *Udanoceratops* and *Montanoceratops*, Kentaro Chiba, Cary Woodruff and Bobby Boessenecker provided assistance with digital modelling and photogrammetry, and Derek Larson provided assistance with Latinization of the genus name. Many thanks to editor Hans-Dieter Sues and reviewers Andy Farke and Brenda Chinnery for constructive comments that improved the manuscript.

ADDITIONAL INFORMATION AND DECLARATIONS

Funding

Funding for this project was provided by an NSERC postdoctoral fellowship, an NSERC L'Óreal-UNESCO for Women in Science fellowship supplement, a National Geographic

Society Waitt Grant, and a Dinosaur Research Institute grant to Victoria Arbour, and an NSERC Discovery Grant to David Evans. The funders had no role in study design, data collection and analysis, decision to publish, or preparation of the manuscript.

Grant Disclosures

The following grant information was disclosed by the authors:

NSERC postdoctoral fellowship, an NSERC L'Óreal-UNESCO for Women in Science fellowship supplement, a National Geographic Society Waitt Grant, and a Dinosaur Research Institute grant to Victoria Arbour, and an NSERC Discovery Grant to David Evans.

Competing Interests

The authors declare that they have no competing interests.

Author Contributions

- Victoria M. Arbour conceived and designed the experiments, performed the experiments, analyzed the data, contributed reagents/materials/analysis tools, prepared figures and/or tables, authored or reviewed drafts of the paper, approved the final draft.
- David C. Evans conceived and designed the experiments, performed the experiments, authored or reviewed drafts of the paper, approved the final draft.

Data Availability

The following information was supplied regarding data availability:

Data is available in the Supplemental Files.

Specimens are accessioned at the following repositories:

Cerasinops hodgskissi, MOR 300, Museum of the Rockies, Bozeman, Montana, USA. *Ferrisaurus sustutensis*, RBCM P900, Royal BC Museum, British Columbia, Canada. *Leptoceratops gracilis*, CMN 8889, CMN 8887, Canadian Museum of Nature, Ottawa,

Ontario, Canada.

Montanoceratops cerorhynchus, MOR 425, Museum of the Rockies, Bozeman, Montana, USA.

Prenoceratops pieganensis, TCM 2003.1.6, TCM 2003.1.2, TCM 2003.1.11, TCM 2003.1.5, TCM 2003.1.1, TCM 2003.1.4, TCM 2003.1.12, TCM 2003.1.3, TCM 2003.1.9, TCM 2003.1.7, TCM 2003.1.8, Children's Museum of Indianapolis, Indianapolis, Indiana, USA.

New Species Registration

The following information was supplied regarding the registration of a newly described species:

Publication LSID: urn:lsid:zoobank.org:pub:D1C60A34-3632-43AD-BCE0-C93D5E26D1B0.

Ferrisaurus LSID: urn:lsid:zoobank.org:act:8430CA06-567E-45A6-B91E-19E51502369E.

Ferrisaurus sustutensis gen. et sp. nov. LSID: urn:lsid:zoobank.org:act:A7F4267C-8CC6-49B6-8E52-2C2148929B14.

Supplemental Information

Supplemental information for this article can be found online at http://dx.doi.org/10.7717/ peerj.7926#supplemental-information.

REFERENCES

- Arbour VM, Evans DC, Simon DJ, Cullen TM, Braman D. Cretaceous flora and fauna of the Sustut Group near the Sustut River, northern British Columbia, Canada. *Canadian Journal of Earth Sciences* (in press).
- Arbour VM, Graves MC. 2008. An ornithischian dinosaur from the Sustut Basin, north-central British Columbia, Canada. *Canadian Journal of Earth Sciences* **45**(4):457–463 DOI 10.1139/E08-009.
- **Brown B. 1914.** *Leptoceratops*, a new genus of Ceratopsia from the Edmonton Cretaceous of Alberta. *Bulletin of the American Museum of Natural History* **33**:567–580.
- Brown CM, Evans DC, Campione NE, O'Brien LJ, Eberth DA. 2013a. Evidence for taphonomic size bias in the Dinosaur Park Formation (Campanian, Alberta), a model Mesozoic terrestrial alluvial-paralic system. *Palaeogeography, Palaeoclimatology, Palaeoecology* 372:108–122 DOI 10.1016/j.palaeo.2012.06.027.
- Brown CM, Evans DC, Ryan MJ, Russell AP. 2013b. New data on the diversity and abundance of small-bodied ornithopods (Dinosauria, Ornithischia) from the Belly River Group (Campanian) of Alberta. *Journal of Vertebrate Paleontology* 33(3):495–520 DOI 10.1080/02724634.2013.746229.
- **Brown B, Schlaikjer EM. 1942.** The skeleton of *Leptoceratops* with the description of a new species. *American Museum Novitates* **1169**:1–15.
- Bustin RM, McKenzie KJ. 1989. Stratigraphy and depositional environments of the Sustut Group, southern Sustut Basin, north central British Columbia. *Bulletin of Canadian Petroleum Geology* 37:210–223.
- **Chinnery B. 2004.** Description of *Prenoceratops pieganensis* gen. et sp. nov. (Dinosauria: Neoceratopsia) from the Two Medicine Formation of Montana. *Journal of Vertebrate Paleontology* **24(3)**:572–590 DOI 10.1671/0272-4634(2004)024[0572:DOPPGE]2.0.CO;2.
- Chinnery BJ, Horner JR. 2007. A new neoceratopsian dinosaur linking North American and Asian taxa. *Journal of Vertebrate Paleontology* 27(3):625–641 DOI 10.1671/0272-4634(2007)27[625:ANNDLN]2.0.CO;2.
- Chinnery BJ, Weishampel DB. 1998. *Montanoceratops cerorhynchus* (Dinosauria: Ceratopsia) and relationships among basal neoceratopsians. *Journal of Vertebrate Paleontology* 18(3):569–585 DOI 10.1080/02724634.1998.10011085.
- **Chinnery-Allgeier BJ, Kirkland JI. 2010.** An update on the paleobiogeography of ceratopsian dinosaurs. In: Ryan MJ, Chinnery-Allgeier BJ, Eberth DA, eds. *New Perspectives on Horned Dinosaurs*. Bloomington: Indiana University Press, 387–404.
- Coombs W. 1971. The Ankylosauridae. PhD thesis. Columbia University, New York, 487.
- Currie PJ, Holmes RB, Ryan MJ, Coy C. 2016. A juvenile chasmosaurine ceratopsid (Dinosauria, Ornithischia) from the Dinosaur Park Formation, Alberta, Canada. *Journal of Vertebrate Paleontology* 26(2):e1048348 DOI 10.1080/02724634.2015.1048348.

- **Dodson P. 2003.** Redescription of neoceratopsian dinosaur *Archaeoceratops* and early evolution of Neoceratopsia. *Acta Palaeontologica Polonica* **48(2)**:261–272.
- Eberth DA, Evans DC, Brinkman DB, Therrien F, Tanke DH, Russell LS. 2013. Dinosaur biostratigraphy of the Edmonton Group (Upper Cretaceous), Alberta, Canada: evidence for climate influence. *Canadian Journal of Earth Sciences* 50(7):701–726 DOI 10.1139/cjes-2012-0185.
- Enkin RJ, Mahoney JB, Baker J, Riesterer J, Haskin ML. 2003. Deciphering shallow paleomagnetic inclinations: 2. Implications from Late Cretaceous strata overlapping the Insular/ Intermontane Superterrane boundary in the southern Canadian Cordillera. *Journal of Geophysical Research: Solid Earth* 108(B4):1–19 DOI 10.1029/2002JB001983.
- **Evans DC, Schott RK, Larson DW, Brown CM, Ryan MJ. 2013.** The oldest North American pachycephalosaurid and the hidden diversity of small-bodied ornithischian dinosaurs. *Nature Communications* **4**:1–10.
- Evenchick CA, Ferri F, Mustard PS, McMechan M, Osadetz KG, Stasiuk L, Wilson NSF, Enkin RJ, Hadlari T, McNicoll VJ. 2003. Recent results and activities of the Integrated Petroleum Resource Potential and Geoscience Studies of the Bowser and Sustut Basins project, British Columbia. *Current Research, Geological Survey of Canada* 13:1–11.
- Farke AA, Maxwell WD, Cifelli RL, Wedel MJ. 2014. A ceratopsian dinosaur from the Lower Cretaceous of western North America, and the biogeography of Neoceratopsia. *PLOS ONE* 9(12):e112055 DOI 10.1371/journal.pone.0112055.
- Fearon JL, Varricchio DJ. 2015. Morphometric analysis of the forelimb and pectoral girdle of the Cretaceous ornithopod dinosaur Oryctodromeus cubicularis and implications for digging. Journal of Vertebrate Paleontology 35(4):e936555 DOI 10.1080/02724634.2014.936555.
- **Fowler DW. 2017.** Revised geochronology, correlation, and dinosaur stratigraphic ranges of the Santonian-Maastrichtian (Late Cretaceous) formations of the Western Interior of North America. *PLOS ONE* **12(11)**:e0188426 DOI 10.1371/journal.pone.0188426.
- Gilmore CW. 1939. Ceratopsian dinosaurs from the Two Medicine formation, Upper Cretaceous of Montana. *Proceedings of the United States National Museum* 87(3066):1–18 DOI 10.5479/si.00963801.87-3066.1.
- **Godefroit P, Lambert O. 2007.** A re-appraisal of *Craspedodon lonzeensis* Dollo, 1883 from the Upper Cretaceous of Belgium: the first record of a neoceratopsian dinosaur in Europe? *Bulletin de l'Institut Royal des sciences naturelles de Belgique, sciences de la terre* 77:83–93.
- Goloboff PA, Farris JS, Nixon KC. 2008. TNT, a free program for phylogenetic analysis. *Cladistics* 24(5):774–786 DOI 10.1111/j.1096-0031.2008.00217.x.
- Han F, Forster CA, Xu X, Clark JM. 2018. Postcranial anatomy of *Yinlong downsi* (Dinosauria: Ceratopsia) from the Upper Jurassic Shishugou Formation of China and the phylogeny of basal ornithischians. *Journal of Systematic Palaeontology* 16(14):1159–1187 DOI 10.1080/14772019.2017.1369185.
- He Y, Makovicky PJ, Wang K, Chen S, Sullivan C, Han F, Xu X, Farke AA. 2015. A new leptoceratopsid (Ornithischia, Ceratopsia) with a unique ischium from the upper Cretaceous of Shandong Province, China. *PLOS ONE* **10**(12):e0144148 DOI 10.1371/journal.pone.0144148.
- **Kobayashi Y, Barsbold R. 2006.** Ornithomimids from the Nemegt Formation of Mongolia. *Journal of the Paleontological Society of Korea* **22**:195–207.
- Lehman TM. 1987. Late Maastrichtian paleoenvironments and dinosaur biogeography in the western interior of North America. *Palaeogeography, Palaeoclimatology, Palaeoecology* 60:189–217 DOI 10.1016/0031-0182(87)90032-0.

- Lehman TM. 2001. Late Cretaceous dinosaur provinciality. In: Tanke DH, Carpenter K, eds. *Mesozoic Vertebrate Life: New Research Inspired by the Paleontology of Philip J. Currie.* Bloomington: Indiana University Press, 310–328.
- Lindgren J, Currie PJ, Siverson M, Rees J, Cederström P, Lindgren F. 2007. The first neoceratopsian dinosaur remains from Europe. *Palaeontology* 50(4):929–937 DOI 10.1111/j.1475-4983.2007.00690.x.
- **Longrich N. 2010.** The function of large eyes in *Protoceratops*: a nocturnal ceratopsian? In: Ryan MJ, Chinnery-Allgeier BJ, Eberth DA, eds. *New Perspectives on Horned Dinosaurs: The Royal Tyrrell Museum Ceratopsian Symposium*. Bloomington: Indiana University Press, 308–327.
- Longrich NR. 2016. A ceratopsian dinosaur from the Late Cretaceous of eastern North America, and implications for dinosaur biogeography. *Cretaceous Research* 57:199–207 DOI 10.1016/j.cretres.2015.08.004.
- Longrich NR, Currie PJ, Zhi-Ming D. 2010. A new oviraptorid (Dinosauria: Theropoda) from the Upper Cretaceous of Bayan Mandahu, Inner Mongolia. *Palaeontology* 53(5):945–960 DOI 10.1111/j.1475-4983.2010.00968.x.
- Maddison WP, Maddison DR. 2011. Mesquite: a modular system for evolutionary analysis. Version 3.04. *Available at http://mesquiteproject.org*.
- Makovicky PJ. 2001. A *Montanoceratops cerorhynchus* (Dinosauria: Ceratopsia) braincase from the Horseshoe Canyon Formation of Alberta. In: Tanke DH, Carpenter K, Skrepnick MW, eds. *Mesozoic Vertebrate Life*. Bloomington: Indiana University Press, 243–262.
- Makovicky PJ. 2010. A redescription of the *Montanoceratops cerorhynchus* holotype, with a review of referred material. In: Ryan MJ, Chinnery-Allgeier BJ, Eberth DA, eds. *New Perspectives on Horned Dinosaurs*. Bloomington: Indiana University Press, 68–82.
- McKenna MC, Love JD. 1970. Local stratigraphic and tectonic significance of *Leptoceratops*, a Cretaceous dinosaur in the Pinyon Conglomerate, northwestern Wyoming. *United States Geological Survey, Professional Paper* **700**:55–61.
- Miyashita TE, Currie PJ, Chinnery-Allgeier BJ. 2010. First basal neoceratopsian from the Oldman Formation (Belly River Group), southern Alberta. In: Ryan MJ, Chinnery-Allgeier BJ, Eberth DA, eds. *New Perspectives on Horned Dinosaurs: The Royal Tyrrell Museum Ceratopsian Symposium*. Bloomington: Indiana University Press, 83–90.
- Morschhauser EM, You H, Li D, Dodson P. 2019. Phylogenetic history of *Auroraceratops rugosus* (Ceratopsia: Ornithischia) from the Lower Cretaceous of Gansu Province, China. *Journal of Vertebrate Paleontology* **38(suppl 1)**:117–147.
- Nessov LA, Kaznyshkina LF, Cherepanov GO. 1989. Ceratopsian dinosaurs and crocodiles of the Mesozoic of Middle Asia. In: Bogdanova TN, Khozatsky LI, eds. *Theoretical and Applied Aspects of Modern Palaeontology*. Leningrad: Nauka, 144–154.
- **Ostrom JH. 1978.** Leptoceratops gracilis from the "Lance" Formation of Wyoming. Journal of Paleontology **52**:697–704.
- **Ott CJ. 2007.** Cranial anatomy and biogeography of the first *Leptoceratops gracilis* (Dinosauria: Ornithischia) specimens from the Hell Creek Formation, southeast Montana. In: Carpenter K, ed. *Horns and Beaks.* Bloomington: Indiana University Press, 213–233.
- Prieto-Márquez A. 2014. A juvenile *Edmontosaurus* from the late Maastrichtian (Cretaceous) of North America: implications for ontogeny and phylogenetic inference in saurolophine dinosaurs. *Cretaceous Research* 50:282–303 DOI 10.1016/j.cretres.2014.05.003.

- Rowe T. 1989. A new species of the theropod dinosaur *Syntarsus* from the Early Jurassic Kayenta Formation of Arizona. *Journal of Vertebrate Paleontology* 9(2):125–136 DOI 10.1080/02724634.1989.10011748.
- Ryan MJ, Currie PJ. 1998. First report of protoceratopsians (Neoceratopsia) from the Late Cretaceous Judith River Group, Alberta, Canada. *Canadian Journal of Earth Sciences* 35(7):820–826 DOI 10.1139/e98-033.
- Ryan MJ, Evans DC. 2005. Ornithischian dinosaurs. In: Currie PJ, Koppelhus EB, eds. *Dinosaur Provincial Park: A Spectacular Ancient Ecosystem Revealed*. Bloomington: Indiana University Press, 312–348.
- Ryan MJ, Evans DC, Currie PJ, Brown CM, Brinkman D. 2012. New leptoceratopsids from the Upper Cretaceous of Alberta, Canada. *Cretaceous Research* 35:69–80 DOI 10.1016/j.cretres.2011.11.018.
- Smith ND, Makovicky PJ, Hammer WR, Currie PJ. 2007. Osteology of *Cryolophosaurus ellioti* (Dinosauria: Theropoda) from the Early Jurassic of Antarctica and implications for early theropod evolution. *Zoological Journal of the Linnean Society* **151(2)**:377–421 DOI 10.1111/j.1096-3642.2007.00325.x.
- **Sternberg CM. 1951.** Complete skeleton of *Leptoceratops gracilis* Brown from the Upper Edmonton Member on Red Deer River, Alberta. *National Museum of Canada Bulletin, Annual Report* **123**:225–255.
- Sues H-D, Averianov A. 2009. *Turanoceratops tardabilis*—the first ceratopsid dinosaur from Asia. *Naturwissenschaften* **96(5)**:645–652 DOI 10.1007/s00114-009-0518-9.
- Tanke DH. 2007. Ceratopsian discoveries and work in Alberta, Canada: Historical review and census. In: Braman DR, ed. Ceratopsian Symposium: Short Papers, Abstracts, and Programs. CD ROM appendix. Drumheller: Royal Tyrrell Museum of Palaeontology, 147–148.
- **Tereschenko VS. 2008.** Adaptive features of protoceratopoids (Ornithischia: Neoceratopsia). *Paleontological Journal* **42(3)**:273–286 DOI 10.1134/S003103010803009X.
- Van Hinsbergen DJJ, De Groot LV, Van Schaik SJ, Spakman W, Bijl PK, Sluijs A, Langereis CG, Brinkuis H. 2015. A paleolatitude calculator for paleoclimate studies. *PLOS ONE* 10(6):e0126946 DOI 10.1371/journal.pone.0126946.
- Zanno LE, Sampson SD. 2005. A new oviraptorosaur (Theropoda, Maniraptora) from the Late Cretaceous (Campanian) of Utah. *Journal of Vertebrate Paleontology* 25(4):897–904 DOI 10.1671/0272-4634(2005)025[0897:ANOTMF]2.0.CO;2.
- Zheng R, Farke A, Kim G. 2011. A photographic atlas of the pes from a hadrosaurine hadrosaurid dinosaur. *PalArch's Journal of Vertebrate Palaeontology* 8:1–12.